

Base from U.S. Geological Survey 7.5' Lake Mathews quadrangle, 1967 Polyconic projection

SCALE 1:24,000 CONTOUR INTERVAL 40 FEET

GEOLOGIC MAP OF THE LAKE MATHEWS 7.5' QUADRANGLE, RIVERSIDE COUNTY, CALIFORNIA

Version 1.0

Douglas M. Morton¹ and F. Harold Weber, Jr.²

Digital preparation by

Van M. Diep¹ and Ursula Edwards-Howells³

¹U.S. Geological Survey **Department of Earth Sciences** University of California Riverside, CA 92521

←13.5°、

² California Division of Mines and Geology 655 South Hope Street Suite 700 Los Angeles, California 90017

Prepared in cooperation with the

1 MILE 1 KILOMETER

Geology mapped by F.H. Weber, 1976;

and D.M. Morton, 1993-5

Department of Earth Sciences University of California Riverside, CA 92521

LOCATION MAP

endorsement by the U.S. Government. atmospheric conditions; therefore, scale and proportions may not be true on plots of this map. Digital files available on World Wide Web at http://geopubs.wr.usgs.gov

morphic foliation			
us foliation			

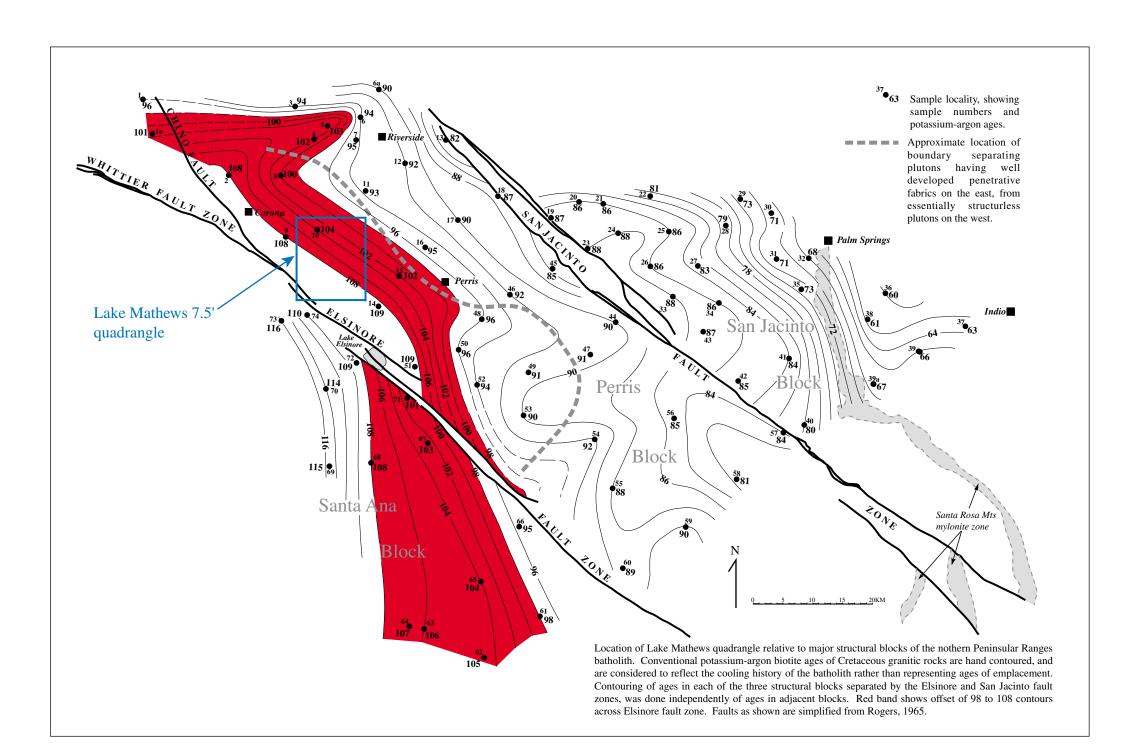
North	Paireotide Paireotide	Pivefist	
South	Lake Mathews 7.5'	Steele Peak	
oPeak	Albertill	Flanore	

This report is preliminary and has not been reviewed for conformity with U.S. Geological Survey editorial standards or with the North American Stratigraphic Code. Any use of trade, firm, or product names in this publication is for descriptive purposes only and does not imply This map was printed on an electronic plotter directly from digital files. Dimensional calibration may vary between electronic plotters and between X and Y directions on the same plotter, and paper may change size due to

DESCRIPTION OF MAP UNITS

	DESCRIPTION OF MAP UNITS
	VERY YOUNG SURFICIAL DEPOSITS—Sediment recently
	transported and deposited in channels and washes, on surfaces of alluvial fans and alluvial plains, and on hillslopes. Soil-profile development is non-
Qaf	existant. Includes: Artificial fill (late Holocene) —Deposits of fill resulting from human construction or mining activities
	YOUNG SURFICIAL DEPOSITS—Sedimentary units that are slightly consolidated to cemented and slightly to moderately dissected. Alluvial fan
	deposits (Qyf series) typically have high coarse:fine clast ratios. Younger surficial units have upper surfaces that are capped by slight to moderately developed pedogenic-soil profiles (A/C to A/AC/B _{cambric} C _{ox} profiles).
Qyf	Includes: Young alluvial fan deposits (Holocene and late Pleistocene) —Gray-hued sand and cobble- and gravel-sand deposits derived from lithicly diverse
	 sedimentary units in the Temescal Valley. In Mockingbird Canyon derived mainly from granitic rock Young axial channel deposits (Holocene and late Pleistocene)—Gray,
Qya	unconsolidated alluvium consisting of fine-grained sand and silt. Occurs in Temescal Valley and channels in dissected very old alluvial fan deposits on the south side of Lake Mathews
Qyv	Young alluvial valley deposits (Holocene and late Pleistocene)—Silty to sandy alluvium on valley floors; gray, unconsolidated
	OLD SURFICIAL DEPOSITS —Sedimentary units that are moderately consolidated and slightly to moderately dissected. Older surficial deposits
	have upper surfaces that are capped by moderately to well-developed
	pedogenic soils (A/AB/B/ C_{OX} profiles and Bt horizons as much as 1 to 2 m thick and maximum hues in the range of 10YR 5/4 and 6/4 through 7.5YR
Qof	6/4 to 4/4 and mature Bt horizons reaching 5YR 5/6). Includes: Old alluvial fan deposits (late to middle Pleistocene)—Indurated, sandy
	and gravely alluvial fan deposits in Temescal Valley. Slightly to moderately dissected; reddish-brown. Some deposits include thin,
0	discontinuous surface layer of Holocene alluvial fan material
Qoa	Old axial channel deposits (late to middle Pleistocene) —Alluvial deposits consisting mainly of sand, but containing minor gravel and silt.
	Gray to reddish-brown, unconsolidated to indurated. Generally slightly dissected. Restricted to several isolated occurrances south of Lake
Qov	Mathews Old alluvial valley deposits (late to middle Pleistocene)—Fluvial
	deposits along valley floors. Consists of moderately indurated, slightlydissected, sandy alluvium, containing lesser silt, and clay-
	bearingalluvium. Some deposits include thin alluvial deposits of Holocene age. Restricted to Gavilan Plateau area in eastern part of
	quadrangle VERY OLD SURFICIAL DEPOSITS—Sediments that are slightly to
	well consolidated to indurated, and moderately to well dissected. Upper surfaces are capped by moderate to well developed pedogenic soils $(A/AB/B/C_{OX})$ profiles having Bt horizons as much as 2 to 3 m thick and
Qvof	maximum hues in the range 7.5YR 6/4 and 4/4 to 2.5YR 5/6) Very old alluvial fan deposits (early Pleistocene)—Mostly well-
	dissected, well-indurated, reddish-brown sand deposits, containing minor gravel. Commonly contains duripans and locally silcretes. Primarily occurs in Temescal Valley and on the south side of Lake
Qvof ₁	Mathews Very old alluvial fan deposits, Unit 1 (early Pleistocene)—Well- dissected, well-indurated, reddish-brown sand deposits, containing minor gravel. Commonly contains duripans and locally silcretes.
Qvoa	Restricted to two small areas on west side of Temescal Valley Very old axial channel deposits (early Pleistocene)—Gravel, sand, and
OTt	silt; reddish-brown, well-indurated, surfaces well-dissected
QTt	Conglomerate of Temescal Wash —Boulder conglomerate, sand and gravel matrix. Fairly well indurated. Brown. Probably deposited on weathered Paleocene surface
TIm	Lake Mathews Formation (Miocene)—Mudstone, conglomerate, and poorly bedded sandstone; massively bedded, nonmarine
Tcgr	Rhyolite-clast conglomerate of Lake Mathews area (Miocene?) —Cobble conglomerate; coarse-grained sandstone matrix, massive bedded,
Тсда	indurated. Cobble clasts include exotic red rhyolite Conglomerate of Arlington Mountain (Paleocene?) —Cobble conglomerate composed of exotic welded tuff clasts with minor clasts
	of exotic quartzite. Found in two small areas north of Arlington
Tsi	Mountain in northwestern part of quadrangle Silverado Formation (Paleocene)—Nonmarine and marine sandstone, and
	siltstone thinly overlying thick basal conglomerate. Basal conglomerate is thoroughly weathered, pale gray to reddish-brown,
	pebble conglomerate, very locally is a boulder conglomerate. Occurs in Temescal Valley
	Rocks of Peninsular Ranges batholith
	Val Verde pluton (Cretaceous)—Relatively uniform pluton composed of biotite-hornblende tonalite. Termed Perris quartz diorite by Dudley
	(1935), Val Verde tonalite by Osborn (1939), and included within
	Bonsall tonalite by Larsen (1948). Name Val Verde adopted by Morton (1999) based on detailed study of Osborn (1939) near Val Verde, a
	former settlement and railway siding midway between Perris and Riverside. Apparently steep-walled Val Verde pluton is eroded to mid-
	pluton level. Emplacement age of pluton 105.7 Ma _{id} . ⁴⁰ Ar/ ³⁹ Ar age of
	hornblende is 100 Ma, biotite 95 Ma and potassium feldspar 88.5 Ma (Snee, pers comm., 1999). Includes:
Kvt	Val Verde tonalite —Gray-weathering, relatively homogenous, massive-to well-foliated, medium- to coarse-grained, hypautomorphic-granular
	biotite-hornblende tonalite; principal rock type of Val Verde pluton.
	Contains subequal biotite and hornblende, quartz and plagioclase. Potassium feldspar generally less than two percent of rock. Where
	present, foliation typically strikes northwest and dips moderately to steeply northeast. Nothern part of pluton contains younger,
	intermittently developed, northeast-striking foliation. In central part of

intermittently developed, northeast-striking foliation. In central part of pluton, tonalite is mostly massive, and contains few segregational masses of mesocratic to melanocratic tonalite. Elliptical- to pancakeshaped, meso- to melonocratic inclusions are common



Gavilan ring complex (Cretaceous)—Composite ring structure consisting of a variety of granitic rocks that range from monzogranite to tonalite. Informally named here for exposures in Gavilan Plateau area, Steele Peak and Lake Mathews 7.5' quadrangles. Western part of complex was termed Estelle quartz diorite and eastern part included in Perris quartz diorite by Dudley (1935). Western part of complex was termed Estelle tonalite and eastern part was included within Bonsall tonlaite by Larsen (1948). Hypersthene is a characteristic mineral of many rocks in complex. Based on texture, depth of erosion is greater in eastern part of compex than in western part. Rocks on west side of the complex commonly have hypabyssal texture and appear to grade into volcanic textrured rock. Several gold mines (e.g., Good Hope, Gavilan, and Santa Rosa mines), which constituted Pinacate mining district (Sampson, 1935), are located within complex. Gold apparently occurrs in arsenopyrite bearing quartz veins. Located in center of ring complex, but not part of it, is near-circular Arroyo del Toro pluton.

Massive textured tonalite-Brown weathering, massive, relatively heterogenous, hypershtene-bearing biotite-hornblende tonalite. Most abundant rock type in complex. Equant-shaped mesocratic to melanocratic inclusions are common. Zircon age is 112.9 Maid and 113.6 Main Foliated tonalite-Gray, medium-grained, foliated biotite-hornblende tonalite containing discoidal mafic inclusions. Most of tonalite lacks

Includes:

Includes

Kgt

Kgtf

Kcg

Kcgd

Kct

 $^{+}_{+}$ K⁺cgq $^{+}_{+}$

Kgb

Khg

Kvspi

Kvem

Kvr

Kvs

Mzu

Mzs

hypersthene. Unit restricted to northern part of complex Cajalco pluton (Cretaceous)-Mostly biotite and biotite-hornblende monzogranite and granodiorite. Informally named here for extensive exposures in Cajalco area, Lake Mathews 7.5' quadrangle. Rocks of Cajalco pluton were included within Cajalco quartz monzonite by Dudley(1935) and within Woodson Mountain granodiorite by Larsen (1948). Unit is a shallow-level pluton emplaced by magmatic stoping within largely volcanic and volcanoclastic rocks. It is tilted eastward and eroded to progressively greater depths from west to east. Upper part of pluton contains a very prominent halo of tourmalinized rock. Zircon ages are 109.5 Ma_{id} and 112.6Ma_{ip} (Premo, pers. comm., 1999).

Tourmalinized monzogranite and granodiorite—Tourmalinized monzogranite and granodiorite that includes some tourmalinized vocanic rock in western part of pluton. Tourmaline is extremely finegrained to aphanitic. Only rock that is essentially all tourmaline is mapped as Kcto. Tourmalinized rock is very resistant to erosion and stands out as small, bold, black hills, locally termed tourmaline 'blow-

Monzogranite-Most of western part of pluton is medium-grained, equigranular, hypautomorphic-granular to subporphyritic monzogranite and subordinate granodiorite. Includes variable amounts of angular inclusions of volcanic rock Granodiorite-Most of eastern part of pluton is medium-grained, equigranular, hypautomorphic granular granodiorite and subordinate monzogranite. Granodiorite includes variable amounts of angular

inclusions **Tonalite**—Masses of mafic biotite-hornblende tonalite. Represents deepest part of pluton Granodiorite and quartz latite, undifferentiated—Nearly equal amounts of putonic and volcanic rock; in some areas, unit is mostly quartz latite. Found near intrusive contacts with Mesozoic volcanic

Granodiorite and gabbro, undifferentiated-Mixed granodiorite and gabbro. In northern and northeastern part of pluton granitic rock contains high concentrations of stoped hornblende gabbro. In some areas granite and gabbro are intimately intermixed producing very heterogneous rock

Generic Cretaceous granitic rocks of the Peninsular Ranges batholith Kgu Granite, undifferentiated (Cretaceous)—Leucocratic fine-to-coarse-

grained massive granite and biotite monzongranite. Most is equigranular and consists of quartz and alkali feldspars. In leucocraitc granite, biotite is a widespread varietal mineral. Muscovite-bearing granite occurs at Bell Mountain, Romoland 7.5' quadrangle Kt Tonalite, undifferentiated (Cretaceous)—Gray, medium-grained biotitehornblende tonalite, typically foliated

Gabbro (Cretaceous)—Mainly hornblende gabbro. Includes Virginia quartz-norite and gabbro of Dudley (1935), and San Marcos gabbro of Larsen (1948). Typically brown-weathering, medium-to very coarsegrained hornblende gabbro; very large poikilitic hornblende crystals are common, and very locally gabbro is pegmatitic. Much of unit is quite heterogeneous in composition and texture Heterogenous granitic rocks (Cretaceous)—A wide variety of heterogenous granititc rocks occur in the Santa Ana Quadrangle. Some heterogneous assemblages include large proportions of schist and gneiss. Rocks in the Santa Ana Mountains include a mixture of

End of rocks of the Peninsular Ranges batholith

rock is the most abundant rock type

monzogranite, granodiorite, tonalite, and gabbro. Tonalite composition

Intrusive rocks associated with Santiago Peak Volcanics (Cretaceous)—Shallow porphyritic intrusive rocks principally of intermediate composition. Composed of plagioclase, clinopyroxene and altered orthopyroxene. Silicic porphyries composed of plagioclase, quartz, and altered pyroxene and biotite (Herzig, 1991) Estelle Mountain volcanics of Herzig (1991) (Cretaceous)—Heterogenous mixture of rhyolite and latite flows, shallow intrusive rocks, and volcaniclastic rocks; andesite is rare. Informally named by Herzig (1991) for exposures in vicinity of Estelle Mountain Rhyolite of Estelle Mountain volcanics of Herzig (1991) (Cretaceous)—Rhyolite; relatively uniform and homogenous Intermixed Estelle Mountain volcanics of Herzig (1991) and

Cretaceous? sedimentary rocks (Cretaceous?)—Complexly intermixed volcanic and sedimentary rocks, which appear to be coeval; sedimentary rocks predominate Mesozoic metasedimentary rocks, undifferentiated (Mesozoic)-Wide variety of low metamorphic grade metamorphic rocks Schist (Mesozoic)-Biotite schist, in part gradational with phyllite. In

lower metamorphic-grade rocks, consists of andalusite-biotite schist. In higher metamorphic-grade rocks, includes cordierite biotite schist, and in highest metamorphic-grade rocks sillimanite schist, and less commonly garnet bearing schist

GEOLOGIC SUMMARY

All but the southeast corner of the Lake Mathews quadrangle is in the Perris block, a relatively stable, rectangular-in-plan area located between the Elsinore and San Jacinto fault zones in the northern Peninsular Ranges Province. In the southwest corner of the quadrangle, a small triangularshaped area that is part of the Santa Ana Mountains structural block, is separated from the Perris block by a short segment of the Elsinore fault zone. The active Elsinore fault zone, a major component of the San Andreas fault system, consists of a series of en echelon northwest-striking right lateral faults located in a graben-like structure.

There is limited relief within the quadrangle because of the presence of two prominent erosion surfaces. The lower Perris surface (about 1,500 feet elevation) has low relief and dominates the physiography of the northern half of the quadrangle. This surface is discontinuously covered by coarsegrained, clastic, non-marine sedimentary rocks of the middle Miocene-age Lake Mathews Formation. A higher Gavilan-Lakeview surface (about 2,100 feet elevation) occurs in the eastern part of the quadrangle, and is locally covered by small exposures of fluvial conglomerate that contain exotic clasts of red rhyolite. The Lake Mathews quadrangle is underlain almost entirely by Cretaceous

and older basement rocks. Two different types of metamorphic rocks are exposed in the quadrangle. In the northeast is a northwest trending exposure of amphibolite grade biotite-bearing schist of probable Mesozoic age. This schist separates massive textured granitic rocks to the west from foliated and layered granitic rocks to the east. The large expanse of metamorphic rock between Temescal Wash and Lake Mathews is low metamorphic grade, typically siliceous, but highly variable in composition. Cretaceous plutonic rocks in the quadrangle are part of the composite Peninsular Ranges batholith, and represent a wide variety of mafic to intermediate composition granitic rocks. Most are massive-textured with the exception of the crudely foliated biotite-hornblende tonalite of the Val Verde pluton in the northeast corner of the quadarangle. The Cajalco pluton, which consists of biotite monzogranite, granodiorite and lesser amounts of biotite-hornblende granodiorite, by far, accounts for most of the granitic rocks in the quadrangle. It is a shallow level pluton emplaced by magmatic stoping into largely intermediate composition volcanic and volcanoclasitic rocks and metamorphic rocks in its western and southern extent and into gabbroic rocks in its northern extent. The pluton appears to be tilted up to the northeast with the texture of the rock changing from subporphyritic rock containing beta-quartz-appearing phenocrysts in the southwestern part of the pluton to coarser-grained hypautomoprhic texture rock in the eastern part. Located in the upper part of the pluton and in overlying wall rock in the shallow western part of the pluton is widespread metasomatic tourmaline rock. Locally parts of the pluton have been completely replaced by tournaline but more commonly tournaline occurs in descrete thin zones, generally along joints. Some of the larger masses of tourmaline rock, locally termed tourmaline 'blowouts', contain cassiterite

and sulfides. One large mass of cassiterite-bearing tourmaline rock supported a tin mining and smelting operation. In the southeast corner of the quadrangle is the northwest part of the Gavilan ring complex. This shallow plutonic complex centered southeast of the quadrangle is predominantly tonalitic composition, characterized by the presence of hypersthene, which is rarely found in Peninsular Ranges batholithic rocks of intermediate composition. Most of the southern part of the quadrangle is underlain by siliceous volcanic and volcanoclastic rocks considered to be coeval with the batholith and which are considered to represent the supra-part of the batholithic magmatism. These rocks generally range in composition from rhyolite to

andesite, but latite is probably the predominant composition. Paleocene continental rocks of the Silverado Formation occur within the Elsinore fault zone and nearby on the adjacent Perris block. Clay-rich parts of the Silveradro Formation have been mined for industrial clay. Near Arlington Mountain, in the northwest part of the quadrangle, are two very small occurrences of conglomerate that consist of exotic welded-tuff clasts and a few exotic bedded quartzite clasts. Extensive Quaternary alluvial deposits are found along the south side of Lake Mathews and in the Temescal Valley along the Elsinore fault zone.

REFERENCES

Dudley, P.H., 1935, Geology of a portion of the Perris block, southern California: California Jour. of Mines and Geology. v. 31, no. 4, p. 487-Herzig, C.T., 1991, Petrogenic and tectonic development of the Santiago

Peak Volcanics, northern Santa Ana Mountains, California; Ph.D dissertation, Riverside, California, University of California, 376 p. Larsen, E.S., 1948, Batholith and associated rocks of Corona, Elsinore, and San Luis Rey quadrangles, southern California: Geol. Soc. of America

Mem. 29, 182 p. Morton, D. M., 1999, Preliminary digital geologic map of the Santa Ana 30' X 60' quadrangle, southern California: U.S. Geological Survey Open-File Report 99-172, 61p., scale, 1:100,000. Rogers, T.H., 1965, Santa Ana sheet: California Division of Mines and Geology Geologic Map of California, scale, 1:250,000. Sampson, R.J., 1935, Mineral resources of a portion of the Perris block Reiverside County, California: California Journal of Mines and Geology, v. 31, p. 507-521.

Streckeisen, A.L., 1973, Plutonic rocks-Classification and nomenclature recommended by the IUGA Subcommission on Systematics of Igneous Rocks: Geotimes, vol. 18, p. 26-30.